Quantitative Field Measurement of Soot Emission from a Large Gas Flare Using Sky-LOSA

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Received July 2, 2010. Revised manuscript received October 28, 2010. Accepted November 18, 2010.

Particulate matter emissions from unconfined sources such as gas flares are extremely difficult to quantify, yet there is a significant need for this measurement capability due to the prevalence and magnitude of gas flaring worldwide. Current estimates for soot emissions from flares are rarely, if ever, based on any form of direct data. A newly developed method to quantify the mass emission rate of soot from flares is demonstrated on a large-scale flare at a gas plant in Uzbekistan, in what is believed to be the first in situ quantitative measurement of soot emission rate from a gas flare under field conditions. The technique, named sky-LOSA, is based on line-of-sight attenuation of skylight through a flare plume coupled with image correlation velocimetry. Monochromatic plume transmissivities were measured using a thermoelectrically cooled scientific-grade CCD camera. Plume velocities were separately calculated using image correlation velocimetry on highspeed movie data. For the flare considered, the mean soot emission rate was determined to be 2.0 g/s at a calculated uncertainty of 33%. This emission rate is approximately equivalent to that of 500 buses driving continuously and equates to approximately 275 trillion particles per second. The environmental impact of large, visibly sooting flares can be quite significant.

Introduction

Emissions of particulate matter (PM) within atmospheric plumes are extremely difficult to quantify. This is especially true for unconfined sources such as gas flares, for which PM emissions in the form of soot are not readily measurable using current techniques. This need persists despite the prevalence of gas flaring in the world, which may exceed 135 billion m³ of gas per year according to recent estimates derived from satellite data (1). Soot emissions from flares are noteworthy not only because of the significant volumes of gas being processed, but also because of the geographic distribution of flares in the world. A large portion of flaring occurs in northern regions of the globe, where the potential for atmospheric transport and deposition of soot aggregates onto the snowpack in Arctic regions is enhanced. This can exacerbate the climate forcing effect of the soot, which when

considered as black carbon is now generally accepted to have a positive warming effect (2) with at least one well-known study suggesting that soot may contribute as much as 55% the effect of CO_2 in the atmosphere (3).

Given the difficulty in quantifying soot emissions from flares, it is not surprising that emission rates or emission factors for these sources are not widely available in the literature. Of the two known studies to have at least partially considered this issue, one reported plume concentration data only (4) and the other concluded simply that soot accounted for "less than 0.5% of the combustion inefficiencies" in the range of test conditions considered (5). Neither work reported direct emission rate data. Most estimates for regulatory purposes are instead based on a very limited set of emissions factors published by the United States Environmental Protection Agency (USEPA) (6). However, recent work (7) has shown that these emissions factors are either based on questionably relevant measurements from predominantly enclosed flares burning landfill gases, or in one case that involves an incorrect interpretation of concentration data as emission rate data.

Measurements of soot emission rates from operating flares under field conditions would thus be highly valuable. However, in the absence of any widely accepted quantitative measurement techniques, regulatory standards in North America are based on estimates of plume opacity by trained human observers, as specified in the USEPA Test Method 9 (ϑ). This approach is essentially an aesthetic measurement, since broadband opacity values cannot be converted to soot emission rates. Furthermore, human-observed opacity measurements are unavoidably subjective and are susceptible to several error sources (ϑ). There have been attempts to modernize Method 9 through the use of digital cameras (10, 11), but these do not address the issue of quantitative soot emission measurement.

Recently, a new technique for quantitative soot emission measurements in flare plumes has been reported (12), which was derived from line-of-sight attenuation (LOSA) techniques commonly used in lab-scaled flames (13–15). Referred to as sky-LOSA, this approach relates monochromatic transmissivity measurements of diffuse sky-light through the plume to soot concentrations via Rayleigh–Debye–Gans Fractal Agglomerate (RDG-FA) theory. Concentration data are combined with a measured plume flow rate to determine a soot emission rate.

The present work reports the successful field application of sky-LOSA to measure the soot emission rate from a large, visibly sooting gas flare. The previous lab-scale approach (*12*) was extended and the underlying theory developed to permit orthogonal interpolation behind an angular plume and to evaluate plume velocity with a high-speed camera using image correlation velocimetry. This setup is the first demonstration of a soot emission measurement with sky-LOSA under field conditions and to the authors' knowledge, this is the first time the soot emission rate in the unconfined atmospheric plume of a flare has been directly quantified.

Theory. The key elements of the theory are summarized here and expanded to derive a formulation for calculating mass emission rate of soot based on measurable parameters in a field experiment. As detailed in (*12*), a measurement of plume transmissivity using sky-LOSA can be linked to the wavelength-dependent extinction coefficient, $K_{\lambda}^{(e)}$, in accordance with the Beer–Lambert–Bouguer Law (*16*), as shown in eq 1:

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$$\tau_{\lambda} = \frac{I_{\lambda}}{I_{\lambda_0}} = \exp[-\int K_{\lambda}^{(e)} \mathrm{d}x] \tag{1}$$

where the transmissivity, τ_{λ} , is measured at a single wavelength, λ , and is defined as the ratio of light intensity after passing through a medium (i.e., a plume), I_{λ} , divided by the intensity of the unobstructed sky-light, I_{λ_0} . The limits of integration are arbitrary as long as they enclose the absorbing medium. I_{λ_0} is obtained by interpolation as discussed in further detail below. Since scattering and absorption of light by soot aggregates can be accurately modeled using RDG-FA theory (17–20), for a soot-laden plume of a flare it is then possible to relate the extinction coefficient in eq 1 to the soot volume fraction. Using RDG-FA theory, the line integrated extinction coefficient along a chord through the plume can be directly related to the line integration of soot volume fraction, f_{v} , along that chord as shown in eq 2:

$$-\ln(\tau_{\lambda}) = \int K_{\lambda}^{(e)} \mathrm{d}x = \frac{1}{\lambda} 6\pi (1 + \rho_{\mathrm{sa}}^*) E(m)_{\lambda} \int f_{\nu} \mathrm{d}x \quad (2)$$

where $E(m)_{\lambda}$ is the soot absorption refractive index function (21-26) and ρ_{sa}^* is the ratio of the light scattering to light absorption by the plume. As discussed in ref 12 it necessary to consider both out-scattering and in-scattering in the calculation of ρ_{sa}^* . For the present analysis, the sky was modeled using the standard clear sky (V.4) of the Commission Internationale de l'Eclairage (CIE) (27) with a sun elevation of 5°. The in- and out-scattering were calculated using RDG-FA theory (17) for the camera angle of 12° from horizontal used in the measurements. A Monte Carlo simulation was used for the calculation of ρ_{sa}^* , for a wide range of soot properties (i.e., primary particle diameters, d_p , from 20 to 50 nm; mean particles per aggregate, N_g from 10 to 300; aggregate distribution widths, σ_g , from 2 to 3.5) and wavelengths from 400–700 nm. The average ρ_{sa}^* is 0.065 \pm 0.065.

Equation 2 can be rearranged and multiplied by soot density and local plume velocity to calculate the local mass flux of soot, perpendicular to a chord through the plume that coincides with the optical axis. As shown in Figure 1, coordinate axes can be defined where x is the direction of the optical axis (perpendicular to the plane of the image), z is the direction of plume propagation, and y is the remaining orthogonal coordinate that defines the "width" of the plume. If the plume velocity is assumed uniform along a chord in the x-direction (along the optical axis), and that chord has an infinitesimal width dy in the y-direction, then the elemental mass flow rate of soot in the z-direction corresponding to this chord can be calculated according to eq 3:

$$d\dot{m}_{soot} = \rho_{soot} u(y) \frac{-\ln[\tau(y)\lambda]}{6\pi(1+\rho_{sa}^*)E(m)_{\lambda}} dy$$
(3)

where is the ρ_{soot} is the soot density, and u(y) is the component of the plume velocity in the *z*-direction. Finally, by integrating across the width of the plume in the *y*-direction, the total mass flow rate of soot in the plume can be calculated, as shown in eq 4.

$$\dot{m}_{\text{soot}} = \frac{-\rho_{\text{soot}}\lambda}{6\pi(1+\rho_{\text{sa}}^*)E(m)_{\lambda}}\int u(y)\ln[\tau_{\lambda}(y)]dy \qquad (4)$$

In practice, we are interested in measuring the mean soot emission rate which is obtained from an ensemble average of N_f measurements using image data collected at different times, as defined in eq 5:



FIGURE 1. Schematic of experimental setup and photograph of the flare used in testing.

$$\overline{\dot{m}_{\text{soot}}} = A \frac{1}{N_f} \sum_{i=1}^{N_f} \int u(y)_i [\ln[\tau_\lambda(y)_i]] dy$$
(5)

where the constant $A = (\rho_{soot}\lambda)/(6\pi(1 + \rho_{sa}^*)E(m)\lambda)$ has been defined for convenience.

The velocity of the plume in the *z*-direction, u(y), can be written in terms of a mean component U(y) and a fluctuating component u'(y) so that

$$u(y) = U(y) + u'(y)$$
 (6)

Substituting eq 6 into eq 5 and rearranging,

$$\overline{\dot{m}_{\text{soot}}} = A \frac{1}{N_f} \sum_{i=1}^{N_f} \int U \ln(\tau_{\lambda})_i \mathrm{d}y + A \frac{1}{N_f} \sum_{i=1}^{N_f} u'_i \ln(\tau_{\lambda})_i \mathrm{d}y$$
(7)

where for simplicity of notation, the variations of each of the parameters, U, u', and τ_{λ} in the *y*-direction (i.e., over the width of the plume) are understood but not explicitly shown. Since the time-averaged value of u' is 0, and assuming that the velocity fluctuations in the plume are associated with mixing and dilution in the plume via atmospheric turbulence but are not correlated with the formation and emission of soot from the flame, then the second term on the right-hand side of eq 7 is reduced to zero.

The mean mass emission rate of soot can be then calculated according to eq 8:

$$\overline{\dot{m}_{\text{soot}}} = A \frac{1}{N_f} \sum_{i=1}^{N_f} \int U(y) \ln(\tau_{\lambda}(y)_i) dy$$
(8)

where dependencies of U and τ_{λ} with y are once again shown explicitly. Equation 8 is particularly useful since it is dependent on the mean velocity profile of the plume, and thus the instantaneous velocity field does not need to be measured synchronously with instantaneous transmissivity. This is the equation that must be solved to quantify the mass emission rate of soot in a plume via a field measurement using sky-LOSA.

Finally, multiple measurements can be made within a single frame at N_z different *z*-locations along the plume and averaged as shown in eq 9 to reduce measurement uncertainty. Furthermore, it is recognized that the order of the summations in this equation can be interchanged, which can simplify processing.

$$\overline{\dot{m}_{\text{soot}}} = \frac{1}{N_f N_z} \sum_{j=1}^{N_z} \sum_{i=1}^{N_f} A \int U(y, z)_j \ln(\tau_\lambda(y, z)_{i,j}) dy \qquad (9)$$

Methodology. Soot emission rates were measured by sky-LOSA for a flare located at a petrochemical plant in Uzbekistan in July 2008. Figure 1 shows a schematic of the apparatus and a photograph of the flare and plume. The stack was 45 m high with a diameter of 1.05 m. Measurements were acquired with cameras mounted on tripods approximately 216 m from the base of the flare. Data were acquired under clear-sky conditions at dusk, which avoided the possibility of direct scattering of sunlight by the plume. In general however, direct sunlight scattering by the plume is a potential error source, since sunlight that is scattered toward the detector will make the plume appear brighter leading to an overestimation of the transmissivity and an underestimation of the soot emission rate. Thus, reported soot emission rates obtained under direct sunlight conditions could be considered conservative.

Velocity Measurement. High-frame rate images (300 frames per second) of the operating flare and plume were acquired using a Casio EX-F1 digital camera equipped with an integral 36-432 mm zoom lens. At this frame rate, the CMOS sensor in the camera was capable of recording data at a spatial resolution of 512×384 pixels. The lens was adjusted so that the spatial resolution in the plume was 55 mm/pixel, based on calibration with the measured stack diameter. A total of 18 300 image frames were recorded (61 s of movie data) while transmissivity data were being separately acquired. Image correlation velocimetry (ICV) analysis was performed on the movie data using DaVis 7.2.1 software (LaVision Inc.), which permitted quantitative measurement of the bulk velocities across the soot plume of the flare.

Plume Transmissivity. Plume images were acquired with a thermoelectrically cooled scientific grade, 16 bit CCD camera (PIXIS 1024BR, Princeton Instruments). A visible light 105 mm lens was mounted to the 1024×1024 pixel CCD to give a spatial resolution in the plume of 23.8 mm/pixel, also calibrated using the measured stack diameter. A 532 ± 3 nm band-pass filter was fixed to the lens to achieve monochromatic measurements. One thousand images were collected with a 50 ms time gate at randomized time intervals ranging between 0.5 and 1.5 s to prevent aliasing. Acquisition was controlled by a customized program written in LabVIEW.

The sky-LOSA processing was implemented in Mathcad (Parametric Technology Corporation) and followed the steps shown via the example images presented in Figure 2. Pixel binning was first applied to the image (image (a)) in order to smooth shot noise effects and reduce the computational expense of image processing. A 4×4 pixel binning was selected leading to a 95.5 mm/pixel resolution in the processed images. A portion of the plume with sufficient free sky area on each side was then selected and rotated to position the plume axis vertically (image (b)). Calculations were performed at multiple cross sections of the plume corresponding to each row of pixels in the rotated images. The sky intensity was interpolated in the region of the plume using the Loess algorithm (28), which creates a weightedpolynomial regression based on the adjacent sky intensity data (image (c)). The optimal interpolation parameters for the Loess algorithm were selected by interpolating sky images without a plume and then comparing the calculated interpolation data to exact sky intensities, as described in ref 12. The optimal sky interpolation was found when the algorithm was performed using sky intensity data from 60×10 binned pixels on each side of the plume; the interpolation span was set to 0.1. Using these parameters, the interpolation algorithm induced a maximum uncertainty of 20% in the measured



FIGURE 2. Sample plume image illustrating the various steps in the processing algorithm to obtain soot emission rate data. (a) Plume image acquired with the 16-bit CCD camera after 4×4 pixel binning. (b) Rotated section of the plume. (c) Rotated image including interpolated sky intensity. (d) Transmissivity image (ratio (b)/(c)). (e) Image of obtained by combination of (d) with ensemble-averaged velocity data. (f) Vertical profile of $U\ln(\tau_{\lambda})dy$.



FIGURE 3. Variations of the integrated optical thickness, $-\int \ln(\tau_{\lambda}) dy$, as a function of interpolation width for three different example cases.

soot emission rate, as summarized in the uncertainty analysis presented below.

Since the plume was unsteady, the location of the plume axis and interpolation region had to be adjusted both along the plume and for each image frame. An automated procedure was developed to first locate the central axis of the local section of plume being evaluated, and then to iteratively determine the optimal interpolation width for analysis. The process was initialized by interpolating sky intensity over a wide portion of sky. The subsequent transmissivity image was used to obtain the plume axis, defined as the central axis of the portion where the transmittance fell below 95%. The sky intensity was then interpolated using different interpolation widths, centered on this axis.

Figure 3 gives examples of the interpolation width optimization analysis. The optical thickness integrated over each horizontal profile within the interpolation area, $-\int \ln(\tau_{\lambda}) dy$, was used as the criterion to optimize the interpolation width. If the interpolated area was too narrow, parts of the plume were excluded (i.e., the sky intensities

variable	reference	value	uncertainty	contribution to uncertainty o <i>f m</i> _{soot}
ρ _{soot} ρ _{sa} E(m)	22, 30–32 evaluated in present work 21–26	1.89 g/mL 0.065 0.334	0.07 g/mL 0.065 0.04	3.7% 6.1% 12.0%
sky interpolation	evaluated in present work following ref 12			20%
и	evaluated in present work	4 m/s		21.3%
spatial calibration	evaluated in present work	95.5 mm/pixel		5%
total				32.7%

TABLE 1. Error Sources for the Soot Emission Rates Evaluated with Sky-LOSA

were not interpolated at the borders of the plume) and the integrated optical thickness was underestimated. The profiles reached a plateau region as the interpolation width was increased, before decreasing when uncertainties associated with larger widths started to dominate. In each of the three cases shown in Figure 3, the optimal interpolation width corresponds to the local maximum within the plateau region on the plots.

Once the sky interpolation was optimized, transmissivity images were obtained via eq 1 (shown by example as image (d) in Figure 2, which is the ratio of the images (b) and (c)). Two-dimensional images of $U\ln(\tau_i)$ within the plume were evaluated by combining the transmissivity images with the velocity data provided by the ICV analysis of the high-speed camera images (image (e)), aligned based on the image of the stack in each data set. Soot emission rates were then obtained by integrating $U \ln(\tau_{\lambda})$ along y and multiplication by the parameter A using soot property data as shown in Table 1. For each image frame, the soot emission profile along the plume axis was calculated (plot (f)) and averaged to provide a quasi-instantaneous measure of the emission rate. Note that in the present case, these measurements are deemed quasi-instantaneous, since we are combining average velocity measurements with instantaneous transmissivity measurements. However, as noted above the final calculation of the mean soot emission rate is unchanged.

A critical advantage of obtaining both velocity and transmissivity data directly from plume images is that it is not necessary that the axis of plume propagation remains perpendicular to the optical axis of the cameras. Any misalignment from $\theta = 90^{\circ}$ simultaneously causes an increase in optical attenuation as the path length through the plume lengthens, which is compensated by a corresponding decrease in the measured plume velocity. For an example case of a locally uniform cross-section plume, these effects both introduce $\cos(\theta)$ terms that exactly cancel. This observation greatly simplifies the implementation of sky-LOSA in a field setting.

Results

Plume Velocity. Figure 4 shows instantaneous and ensembleaveraged velocity fields in the plume as determined from image correlation velocimetry of the high-speed digital movie images. The vector plots are superimposed on the same example grayscale camera image, which corresponds to the plotted instantaneous velocities. An adaptive mask was used during processing to exclude flame regions as well as regions outside the plume image. Although the original images were captured at 300 frames per second, it was determined during processing that for the calibrated spatial resolution in the plume of 55 mm/pixel, better results were achieved by processing at 60 frames per second so that interframe displacements could be more accurately resolved. As shown in Figure 4b, the ensemble-averaged velocity vectors were quite uniform across the plume, with magnitudes mostly between 4 and 5 m/s and showing some slight divergence as the flow moves downstream.

Although it is not possible to directly evaluate the streamwise velocity variation along the optical axis, the relatively minor variation across the plume in the image plane implies that it would not be significant. Close inspection of the instantaneous images also suggests the same, since the expected 2D shear layer structures on the top of the plume are clearly visible and do not appear smeared due to variation in velocity along the optical axis. These observations are consistent with detailed measurements on nonreacting jets in crossflow (29), which show that velocity variation along the *x*-axis falls to less than 10% of the wind speed by 12 diameters downstream of the jet origin. Finally the



FIGURE 4. Example instantaneous (a) and ensemble average (b) velocity fields calculated from the high-speed movie data.



FIGURE 5. (a) Soot emission rates calculated for each of 200 consecutive frames (black dots). The gray squares represent mean soot emission values for 25 consecutive frames. (b) Average soot emission profiles along the plume axis. The error bars account for frame to frame variations as well as interpolation uncertainty.

transmissivity of the plume is such that the camera is able to "see" along chords through the plume producing a concentration weighted velocity in which the soot is the tracer for the image correlation measurements. This is ideal, since it is the concentration weighted mean-velocity along the optical axis that is desired for the mass flux calculation.

Soot Emission. The frame-averaged soot emission rates are plotted for 200 consecutive acquisitions (dots) in Figure 5a. Also shown are 25-frame running averages (gray box symbols), which range from approximately 1.75 to 2.25 g/s and are centered about the average soot emission rate of 2.0 g/s. The frame-to-frame standard deviation was 0.38 g/s, or approximately 19% of the averaged value. However, the apparent fluctuations among individual frames are primarily an artifact of the calculation method, in which mean velocity data were combined with instantaneous LOSA data for each frame, as explained in the theory presented above. If it were possible to simultaneously and accurately measure synchronized velocity and transmissivity data on a single camera, then these fluctuations would be significantly reduced.

However, as noted above, so long as the fluctuating component of the velocity and the instantaneous transmissivity along profiles through the plume are randomly correlated, then these apparent fluctuations can be accurately removed via averaging. Since currently available CCD- or CMOS-based camera technologies necessitate a trade-off between high frame rates and image noise and sensitivity, the present approach allows hardware requirements to be easily split among two devices. Finally, the overall horizontal trend in the 25-frame averaged data implies that flow conditions were reasonably steady during measurements.

Soot emission rates averaged over 200 consecutive acquisitions are plotted as a function of distance along the plume axis in Figure 5b. The variation with distance is small relative to the uncertainty of the measurements, demonstrating that the diagnostic is relatively insensitive to the measurement location.

Discussion

As described in ref (12), the uncertainty of the soot optical properties is a major source of error for sky-LOSA. As shown in Table 1, the uncertainty of E(m) on its own leads to a 12% uncertainty in the soot emission rate. The associated uncertainty will only be reduced by improving knowledge of the soot optical properties.

The sky interpolation process also generates uncertainty since the algorithm does not give access to the exact sky intensity, as explained in the Methodology. Through comparison of the interpolated image to the exact image, the nominal interpolation uncertainty was quantified at 20%. This high uncertainty was due to speed limitations of the camera's integrated image plane iris-shutter, resulting in a "hot spot" in the middle of the images that was difficult to accurately interpolate. By operating at longer exposure times or by modifying the camera to replace the included shutter with a focal plane shutter, it should be possible to reduce this uncertainty in future measurements.

The uncertainty of the ICV processing to calculate plume velocities is a more significant source of error in the present study. The velocity uncertainty can be conservatively estimated based on the spatial resolution of the images, the calibration scale, the effective exposure time of the camera, the probability distribution of instantaneous vector displacements, and the time interval between frames. For the relatively coarse resolution high-speed video images available in the current study, the spatial calibration uncertainty is estimated at 5.3%. The uncertainty associated with movement of the plume during a frame can be estimated at no more than 20% assuming a worst-case maximum exposure duration of 1/300 s relative to the frame rate of 1/60 s used in processing. Finally, analysis of histograms of the instantaneous velocity results suggests that the ICV correlation peaks are well-defined such that interframe plume displacements as small as 0.05 pixels can be resolved, leading to an estimated absolute uncertainty of 0.2 m/s. For a typical velocity in the plume of 4.0 m/s, these errors can be combined to give estimated uncertainties of approximately 21.3%, which directly affects the uncertainty in the measured soot emission rate as summarized in Table 1.

As shown in Table 1, the total estimated uncertainty on the measured soot emission rates is 33%. While this is certainly non-negligible, put into the perspective of the absence of currently available quantitative measurement techniques for unconfined sources such as flares, this level of accuracy is quite reasonable. By contrast, measurements based on simple opacity as per EPA Method 9 do not offer any quantitative information about soot emission rates and can be prone to significant errors even as a qualitative method (9). Thus, a 33% uncertainty is a reasonable value for a first field demonstration of sky-LOSA and represents a significant improvement over current standards. Moreover, future improvements in the sky-LOSA procedure should permit uncertainties to be reduced. Specifically, recent developments in high-speed sCMOS cameras (33) should very soon enable better picture dynamic range and high frame-rates that would allow the complete sky-LOSA analysis using a single camera. Instantaneous concentration and velocity images would then be acquired simultaneously and with improved spatial resolution, thus significantly reducing the uncertainties in measured velocities and removing the requirement to make calculations using ensemble-averaged velocity data.

Since composition and flow rate of the gas being directed to the flare were not known, it was not possible to calculate a relative soot emission factor from the present data. However, the calculated soot emission rate of 7400 g/h can be compared to typical emission values for vehicles to provide some context. Keogh et al. (34) presented average PM2.5 emission rates of 299 mg/km for typical buses measured during on-road testing. This corresponds to 15 g/h for buses driving at an average speed of 50 km/h. The measured flare emission rate is then equivalent to soot emissions of \sim 500 buses constantly driving. Assuming present soot aggregates have an average composition of 120 primary particles with mean diameter of 39.4 nm (values for postflame soot measured in an inverted methane/air flame (35)) and density of 1.89 g/mL from Table 1, this would correspond to an emission rate on the order of 275 trillion soot aggregates per second. Both statistics indicate the potentially dramatic environmental impact of gas flaring.

Acknowledgments

This work was supported by Natural Resources Canada (project manager, Michael Layer), the Canadian Association of Petroleum Producers (CAPP), and the World Bank Global Gas Flaring Reduction (GGFR) partnership. We gratefully acknowledge Brian Crosland (Carleton University) for help with preparations and David Picard (Clearstone Engineering) for assistance with field measurements in Uzbekistan.

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ES102230Y